

# Optimal Initial Conditions for Coupling Ice sheet Models to Earth System Models

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Motivation

Methods

Synthetic & Realistic Applications

Summary & Future Work





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# Uncertainty Quantification

Uncertainty in predictions from ice sheet models (e.g. sea-level rise) come from multiple sources, including predominately:

- (1) Uncertainties in model forcing - largely related to uncertainties in future climate, which are being explored through emissions-scenario-dependent and perturbed physics ensemble analyses \*\*
- (2) Model uncertainties – largely due to uncertainties in initial and boundary conditions \*\*

PISCEES UQ will focus on the latter, specifically:

- (i) Constraining uncertain initial and boundary condition parameters**
- (ii) Estimating parameter uncertainties using a combination of intrusive (adjoint) and non-intrusive (sampling) approaches \*\***
- (iii) Forward propagation of input parameter uncertainties to assign uncertainties to ice sheet model outputs of interest

*\*\* See talks by: Heimbach et al., Jackson et al. \*\*  
& poster by Salinger et al. (#43, room 1)*

# Motivation

Existing ice sheet model initialization methods do not couple smoothly with realistic climate forcing (from models or observations).

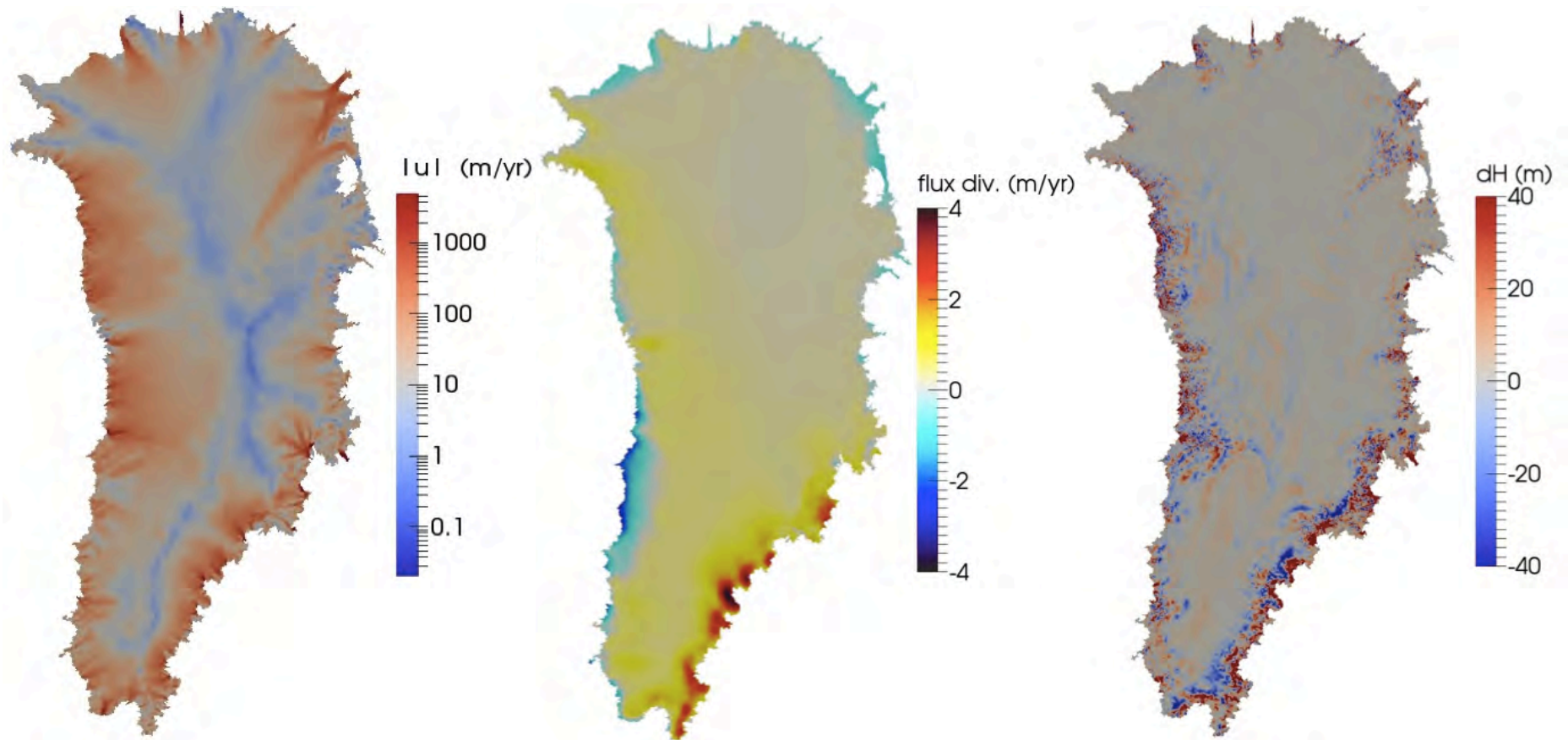
- 1. Spin-up:** initial condition is consistent with climate forcing & long-term transients, but difficult (impossible?) to combine with the goal of closely matching present-day obs. (geometry, flux, etc.)
- 2. Optimization:** fix model geom. to obs., tune model parameters (e.g., basal sliding coeff.) to provide optimal match to observed vels.

Method (1) is impractical because (i) ice dynamic response on  $10^1$ - $10^2$  yr timescales is very strong function of initial geom. and vel., and (ii) spin-up of appropriate duration ( $10^4$ - $10^5$  yr) not practical for high-res., next-gen. ice sheet models.

Method (2) provides good match to present-day obs. but generally leads to unphysical “shock” when coupling to SMB from climate model.



# Motivation



**At equilibrium, the SMB is balanced by the flux divergence**

SMB = surface mass balance = ice accumulation less melting & sublimation

A photograph of a large iceberg in the ocean. The iceberg is white and has a jagged, mountain-like shape. It is surrounded by smaller icebergs and icebergs in the foreground. The water is a deep blue color. The sky is a pale blue color.

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# Optimization Problem

To avoid non-physical shocks when coupling ice sheet models to ESMs, the model flux divergence must be balanced by the model SMB.

$$\frac{\partial H}{\partial t} = -\underbrace{\text{div}(\mathbf{U}H)}_{\text{flux divergence}} + \tau_s, \quad \mathbf{U} = \frac{1}{H} \int_z \mathbf{u} dz \quad \text{At equilibrium: } \text{div}(\mathbf{U}H) = \tau_s$$

surface mass balance

$$\text{Basal boundary condition: } (\sigma \mathbf{n} + \beta \mathbf{u})_{\parallel} = \mathbf{0} \quad \text{on } \Gamma_{\beta}$$

At the same time, we want the initial model geometry and velocity field to match present-day observations (obtained by optim. “beta” field)

**Solution:** PDE-constrained optimization with constraints on both velocity (commonly applied) *and* flux divergence (novel), *additionally* accounting for uncertainties in ice sheet geometry (thickness)

**Optimization Problem:** Find  $\beta$  and  $H$  that minimize the cost functional  $\mathcal{J}$ :



# Optimization Problem

Find  $\beta$ ,  $H$  that minimize the objective functional

$$\begin{aligned}
 \mathcal{J}(\beta, H) = & \int_{\Sigma} \frac{1}{2\sigma_u^2} |\mathbf{u} - \mathbf{u}^{\text{obs}}|^2 ds && \text{surface velocity mismatch} \\
 & + \int_{\Sigma} \frac{1}{2\sigma_{\tau}^2} |\nabla \cdot (\mathbf{U}H) - \tau_s|^2 ds && \text{SMB mismatch} \\
 & + \int_{\Sigma} \frac{1}{2\sigma_H^2} |H - H^{\text{obs}}|^2 ds && \text{thickness mismatch} \\
 & + \mathcal{R}(\beta, H) && \text{Regularization terms}
 \end{aligned}
 \left. \vphantom{\int_{\Sigma}} \right\} \begin{array}{l} \text{Common} \\ \text{Novel} \end{array}$$

**subject to ice-sheet model equations**

(high-order approximation of nonlinear Stokes equations).

# Numerical / Computational Details

## Fwd model (PDE constraint):

- 1<sup>st</sup>-order Stokes approximation<sup>1</sup> (FELIX prototype model)
- FEM discretization
- variable resolution, tetrahedral mesh (min. res. ~4 km)

## Numerical method:

- Quasi-Newton using LBFGS for cost function minimization
- cost function gradients provided by fwd model adjoint <sup>1</sup>

## Software Frameworks:

- *LifeV* FEM library
- *Trilinos*:
  - NOX – Newton nonlinear solver in fwd model
  - AztecOO – PCG linear solver in fwd model
  - ROL – Rapid Optimization Library (ROL) for LBFGS

<sup>1</sup> Perego et al. (2012)





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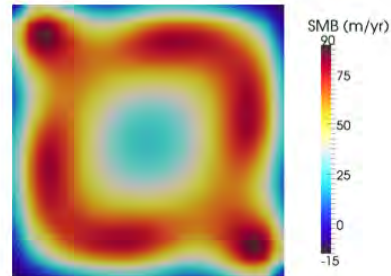
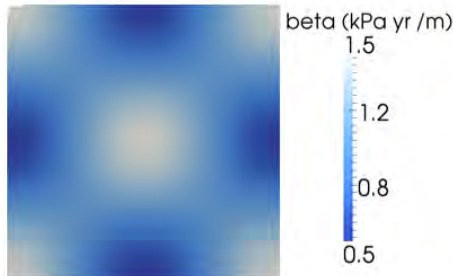
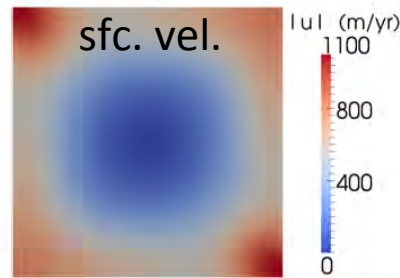
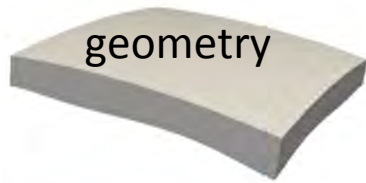
**Synthetic & Realistic Applications**

Summary & Future Work



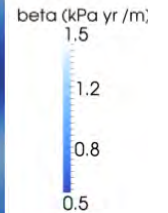
# Synthetic Application

## Problem Setup

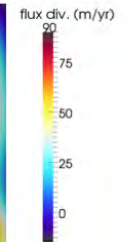
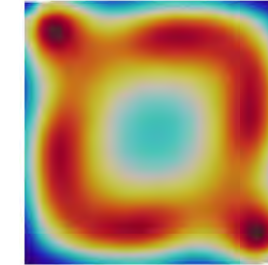
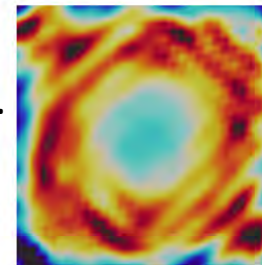


## Inversion Results

sliding  
coeff.



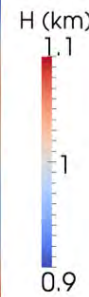
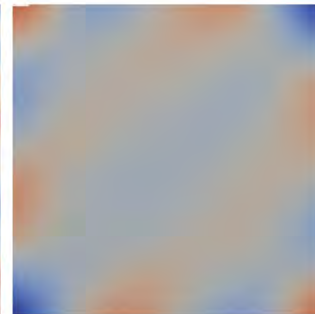
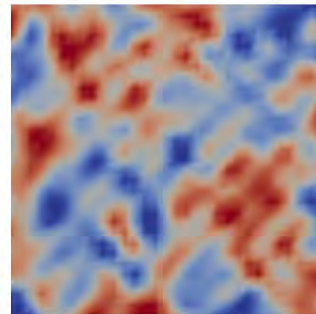
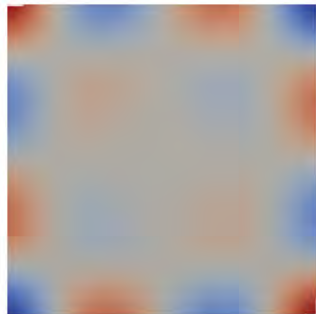
equilib.  
SMB



common

novel

## Thickness



truth

obs. + error

recovered

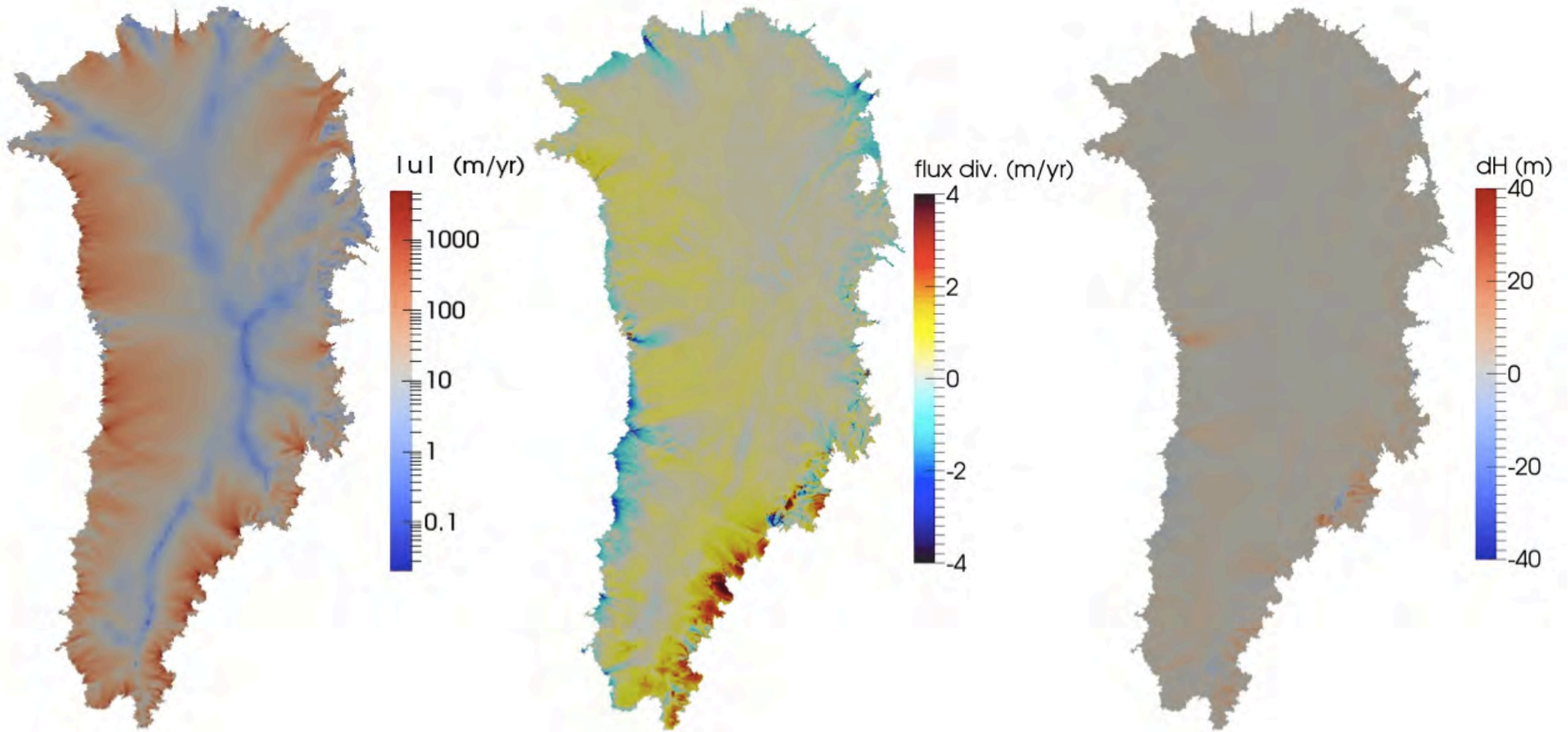
- Common and novel method succeed in recovering sfc. vel.
- Common method does not allow for equilib. SMB constraint
- Common method has difficulty with uncertain thickness data
- Novel method addresses both of these failings

## error in recovered velocity





# Realistic Application

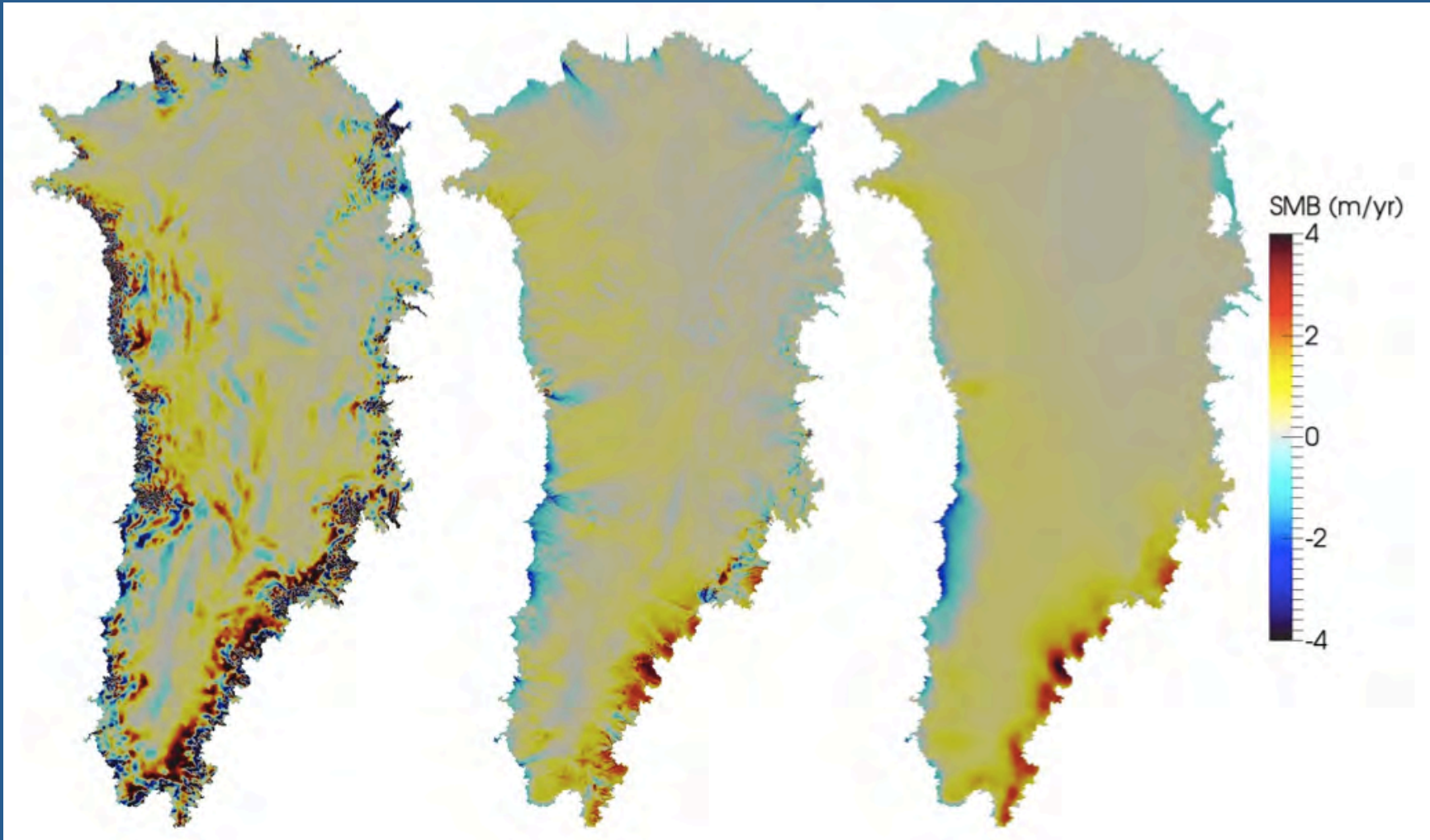


# Realistic Application

Flux Divergence  
(standard optim.)

Flux Divergence  
(improved optim.)

Target SMB







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# Summary & Future Work

Demonstration of promising optimization framework for generating realistic ice sheet model initial conditions that also couple smoothly to climate models

... but how does this relate to UQ?

## Next Steps:

Use approximate Hessian to aid in sampling of posterior parameter distributions \*\* (in addition to finding the MAP point, as shown here)

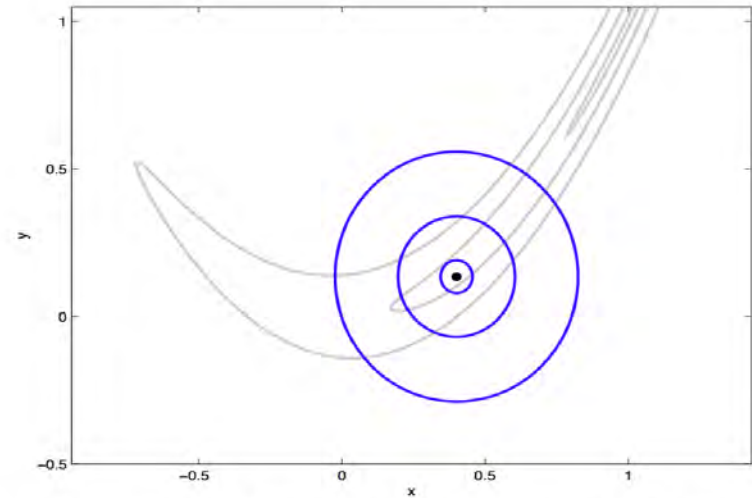
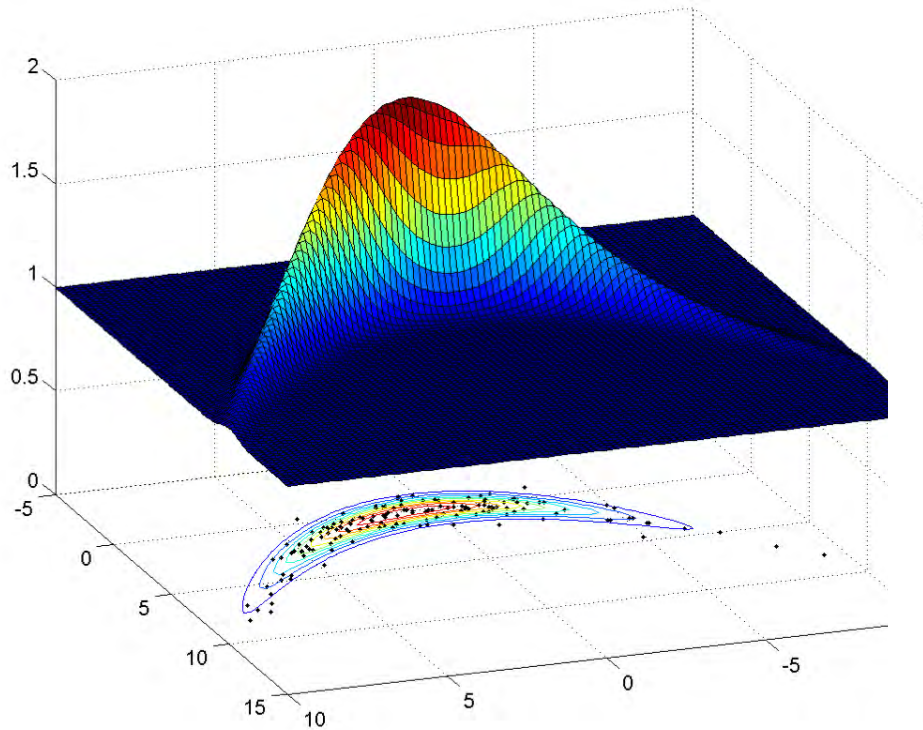
Conduct fwd simulations from these approximate distributions in order to assess resulting uncertainties on model outputs of interest (e.g. ice sheet mass loss and sea-level rise)

\*\* Requires a combination of intrusive (adjoint) and non-intrusive (sampling) methods, including human and software support for both

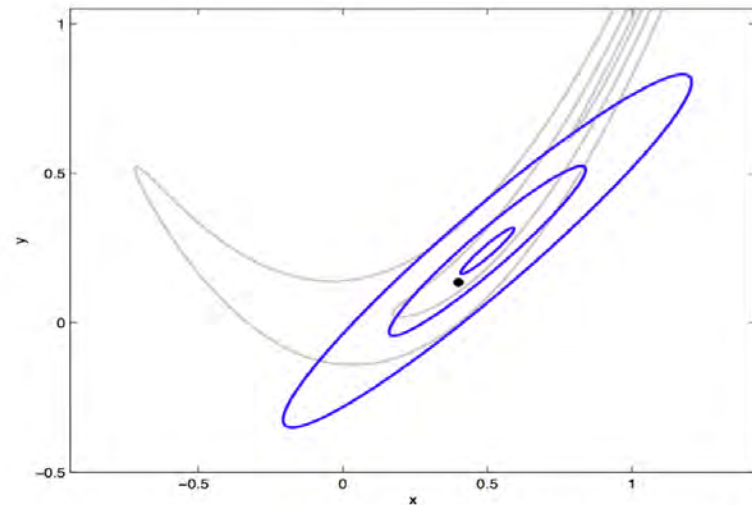


# Future Work

Use of approximate Hessian to increase acceptance rate of MCMC sampling of posterior parameter distribution



The contours of the random walk proposal function overlaid on the Rosenbrock contours.



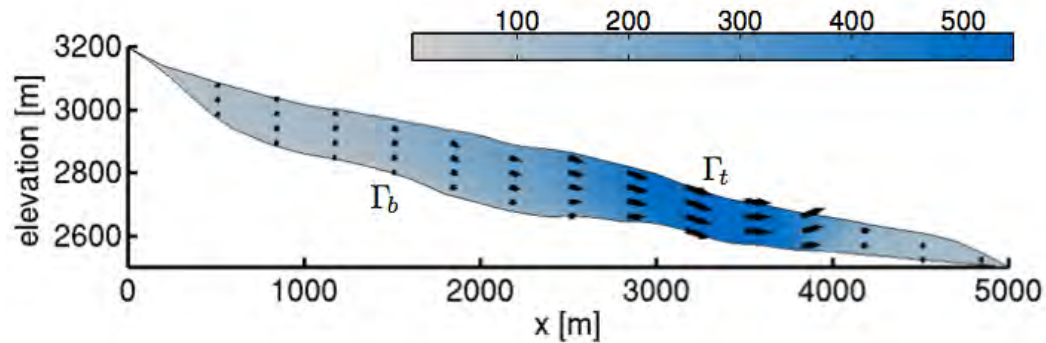
The contours of the stochastic Newton method proposal function overlaid on the Rosenbrock contours.

Figures courtesy of G. Stadler

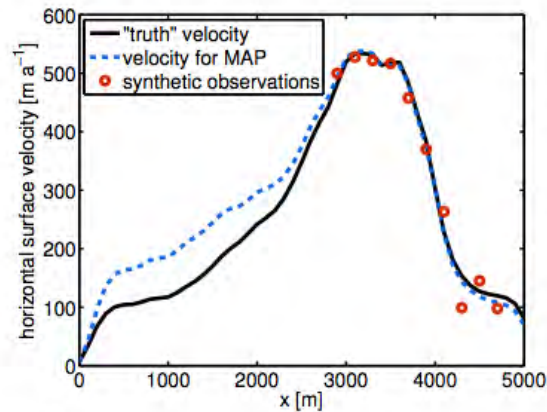


# Future Work

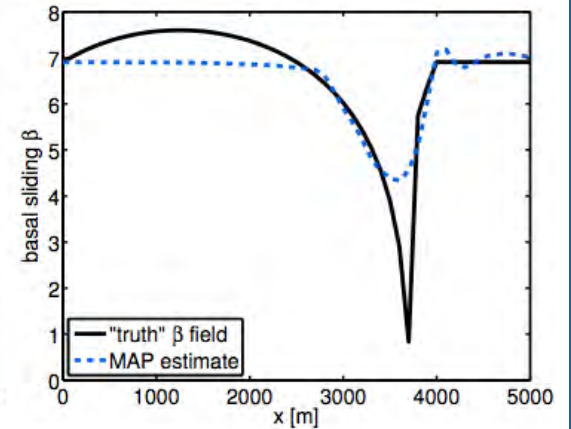
Modeled 2d (x,z) velocity field with basal sliding coefficients tuned to match observations



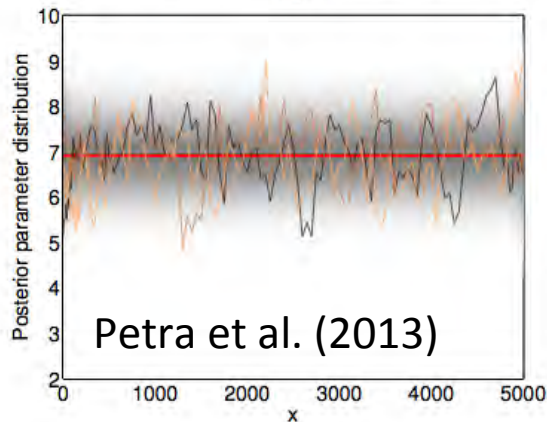
**Left:** modeled (blue) and “true” (blk) surface velocity profile and synthetic observations (red)



**Right:** MAP estimate (blue) and “true” basal sliding coefficient (blk)



**Left:** prior estimate for basal sliding coefficient distribution



**Right:** posterior coefficient distribution obtained using Hessian-informed MCMC sampling

